



RSM2010,



Effects of freshwater forcing on a simulated climatology in the HadGEM2

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Introduction

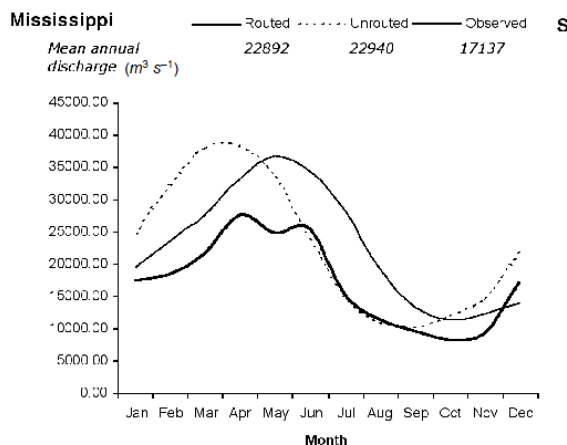


Lucas-Picher et al. 2003; Liston et al. 1994; Oki et al. 1999

- They have been done to determine the **quantity of flow along river** at various points in time and space.
- They focused on **small watersheds** using the LSM or RCM.

Oki et al. 1999

- They **calculated the mean runoff** estimated by the LSM for drainage areas upstream of 250 operational gauging stations.



Mean Annual Runoff [mm/y] by Mean 11 models for 1988

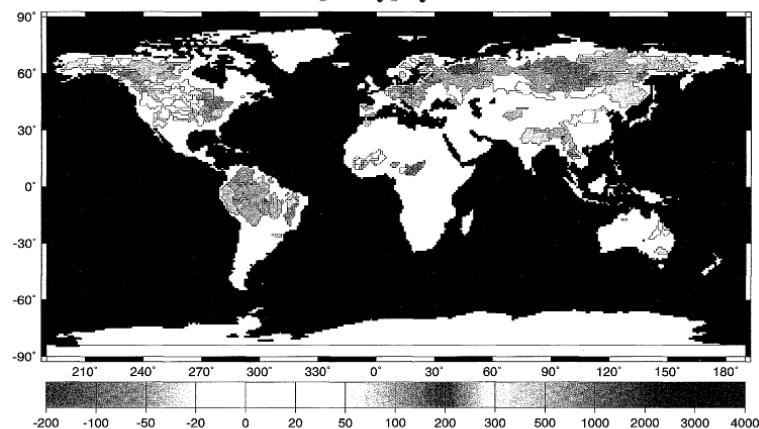


Fig. 4. Mean annual runoff [mm/y] by 11 LSMs averaged within each incremental drainage area for 1988.

Lucas-Picher et al. 1994

- They showed that the **inclusion of flow routing in the Canadian RCM** produces a better agreement with observation-based streamflow estimates.

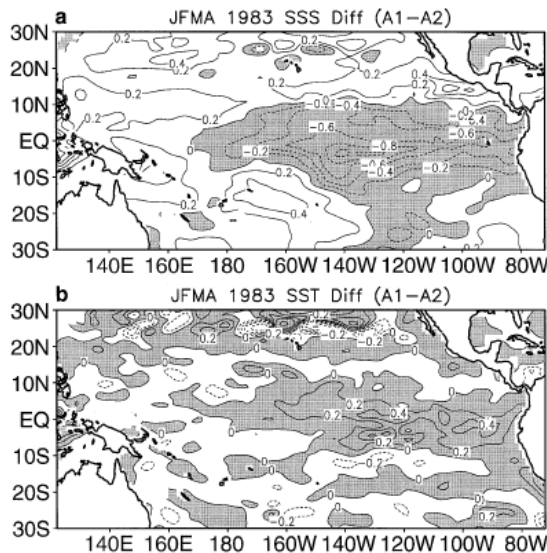


Introduction



Yang et al. 1999; Zhang and Busalacchi 2009; Wu et al. 2010

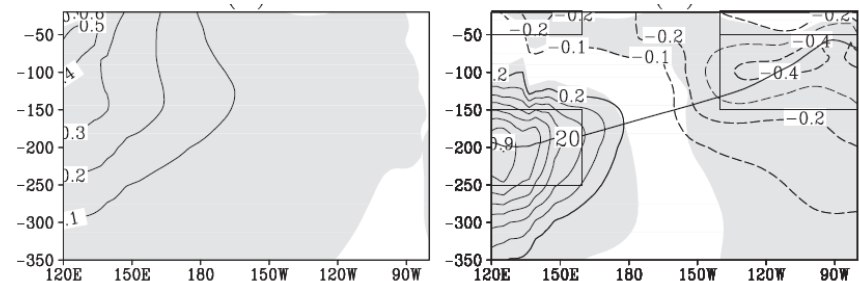
- Climate modeling community has focused on the **need for routing models** to track the flow of water from continents to oceans **at global scale**.
- The simulated freshwater flux into the oceans alters their **salinity** and may affect the **thermohaline circulation**.



Yang et al. 1999

- They showed that the **inclusion of freshwater input from precipitation** could lead to increases in SST by as much as 0.5 K in the tropical ocean **in an OGCM**

Wu et al. 2010



- Using the **fully coupled ocean-atmosphere GCM**, the mechanisms response to **idealized freshwater flux forcing** in the western tropical Pacific were investigated.



Introduction



In this study,

fully coupled GCM + real freshwater forcing

the effects of river flows in the HadGEM2-AO

on a simulated climate are investigated.

HadGEM2-AO version 6.6

Sensitivity test

🔧 Initial time : 0000UTC 1 Sep, 1978

🔧 Analysis : 1979 – 1988 (10yr)

🔧 Resolution:

Atmosphere: 1.25°X 1.875°(horizontal); 38 levels (vertical)

Ocean: near 1°X 1°; 360*216 (horizontal); 40 levels (vertical)

🔧 River Routing scheme (TRIP; Oki and Sud 1998)

🔧 Land surface scheme (MOSES; Cox et al. 1999)

🔧 RIV – Including River Routing scheme

🔧 nRIV – Not including River Routing scheme

Features of MOSES

- ❖ The **Met Office Surface Exchange Scheme** (MOSES), **Cox et al.1999**
- ❖ **4 layer soil hydrology model**
based on a discretised version of the Richards' equation
- ❖ Including of the effects of **soil water phase changes**
on the permeability and heat balance of the soil
- ❖ The surface energy partitioning:
Pennman-Monteith flux equations using a resistance formulation
- ❖ **The sensible heat and the various evaporation** components:
dependent on atmospheric stability and the surface roughness
- ❖ **Evapotranspiration** from soil moisture store:
limited by a surface resistance that is represented
by a “bulk stomatal” resistance model

→ hydrology section

: calculates **the interception of rainfall** by the plant canopy, and the “surface” and “drainage” components of the runoff.



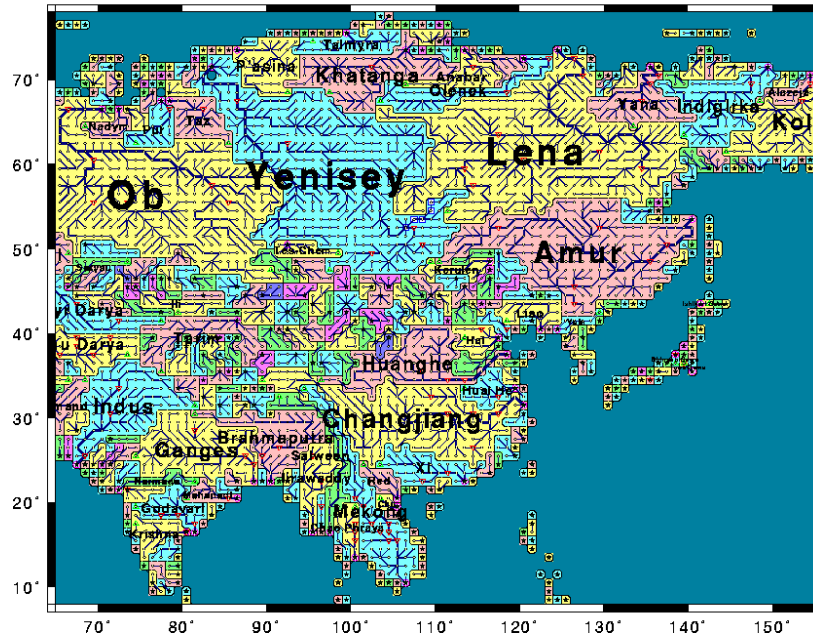
TRIP (Total Runoff Integrating Pathways)



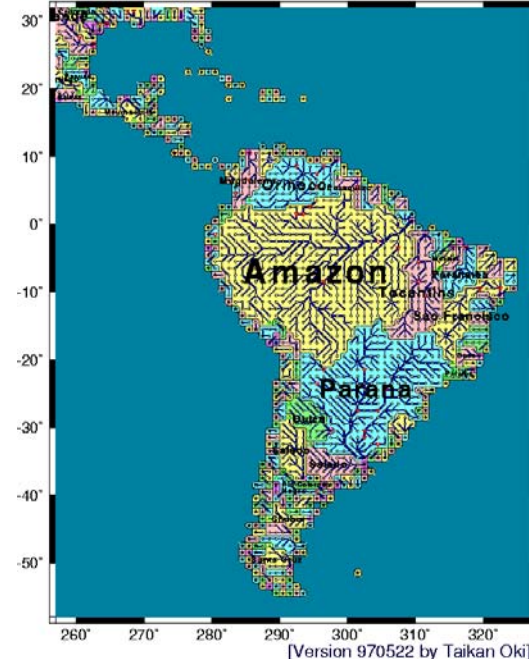
What is TRIP ?

- developed by **Oki and Sud (1998)**
- homepage <http://hydro.iis.u-tokyo.ac.jp/~taikan/TRIPDATA/TRIPDATA.html>
- advection method to route total runoff along prescribed river channels
- representation of isolate the river basins, inter-basin translation of water through river channels, as well as collect and route runoff to the river mouth(s) for all the major rivers.
- To produce the first guess of TRIP, river flow directions were determined from a global DEM (ETOPO5, Edwards 1986).

Rivers in Asia on TRIP in 1°x1° mesh



Rivers in South America on TRIP in 1°x1° mesh

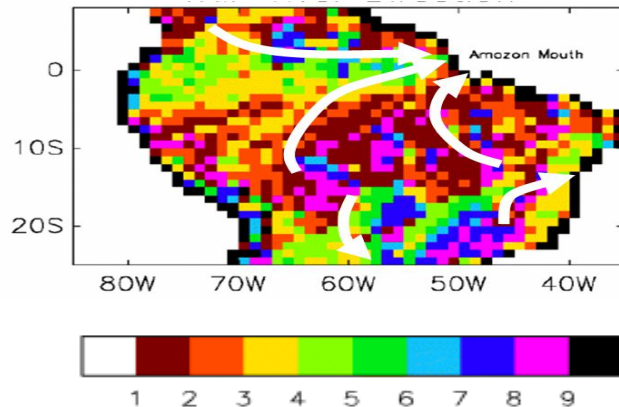




River channel in TRIP

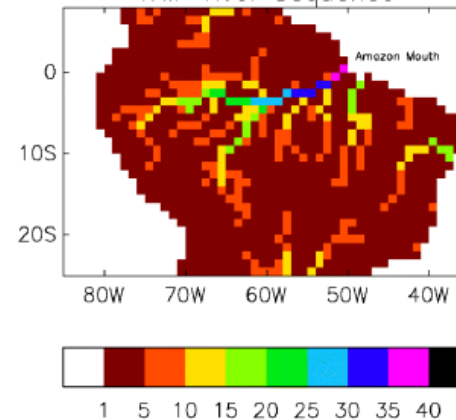


1) Direction



1:N, 2:NE, 3:E, 4:SE, 5:S, 6:SW,
7:W, 8:NW, 9:river mouth
(1 to 8 clockwise from north)

2) Sequence



Each box in the channel - numbered starting at 1 (head of a channel) until it meets a major channel or reach the sea

Features of TRIP

1. The outflow direction was toward the **lowest land point** of the eight neighboring grids, provided the point was lower than the originating point.
2. If a grid point was lower than all neighboring grids, it was marked as "hollow" and was left for the subjective evaluation and **manual correction**.
3. **No** river channels are allowed to **crisscross**.
4. All river channels flow **from one land grid box to another**.
5. Every land grid box has one, and **only one river mouth** toward its down-stream, which eliminates the possibility of grid boxes counterflowing toward each other



Linear routing algorithm of TRIP



Conservation of river storage

$$\frac{dS_{rc}}{dt} = D_{IN} - D_{OUT}$$

S_{rc} : river water storage (kg) in a gridbox

D_{IN} : sum of inflow (kg/s) from neighboring gridboxes runoff (kg/s) produced within the gridbox

D_{OUT} : outflow from a gridbox (one direction only),

~ cS_{rc} , c = u/d, u : effective velocity (m/s) (Miller et al. 1994)

d : distance (m) across the gridbox * meandering ratio



Meandering ratio = a : b

input: runoff(surface + subsurface) from LSM

output: daily mean outflow at river mouths

output (1 * 1) → regridded between the
HadGEM2 and TRIP grids

→ providing output to ocean grid
(Falloon et al., 2007)

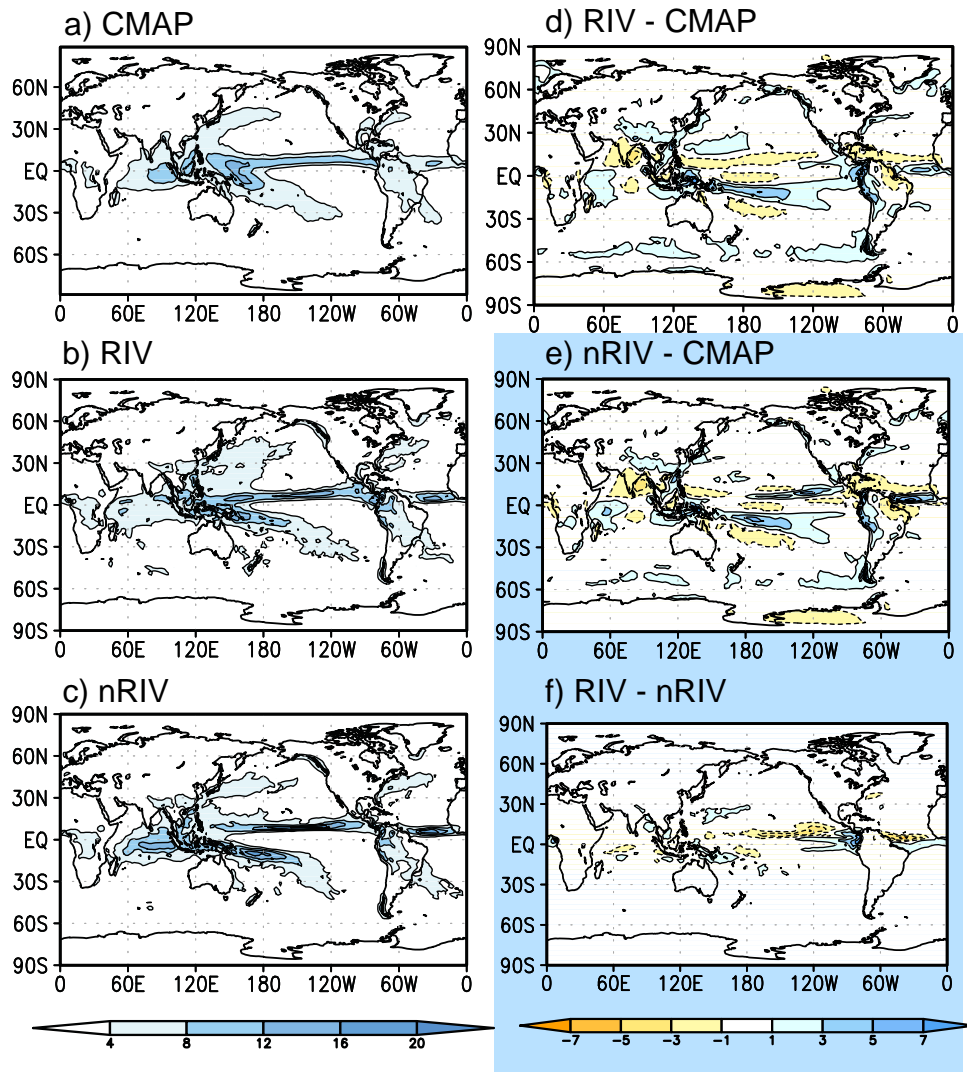


**Results ~
(10yr mean)**



Effects of river routing processes

Precipitation



nRIV & RIV : satisfactorily reproduced.

nRIV run

overestimated:

(central tropical ocean, western South America, central Indian Ocean)

underestimated:

(eastern equatorial Pacific, Atlantic Ocean)

RIV run

decreased

over the **north of the equator** in the central Pacific,

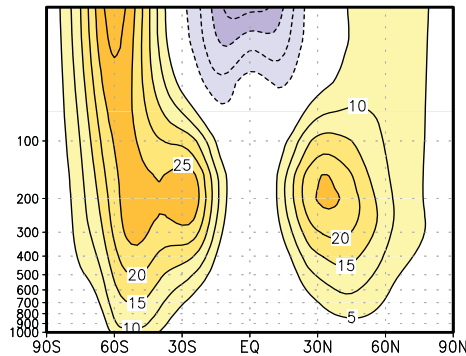
increased

over the **eastern equatorial Pacific**

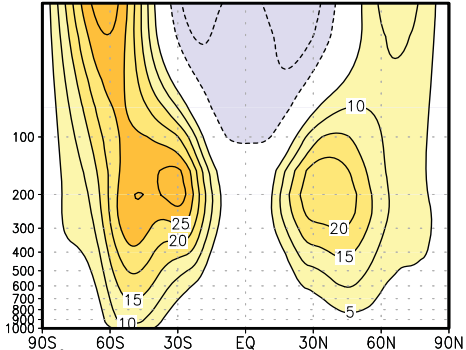


Effects of river routing processes

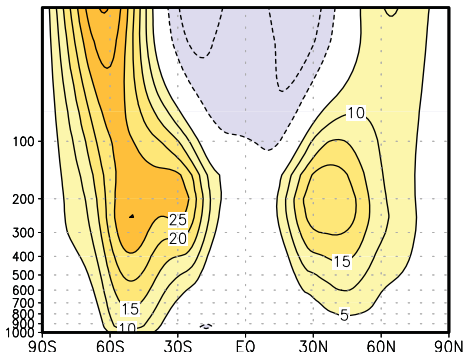
a) R2



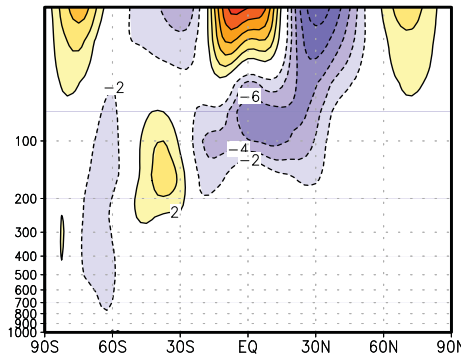
b) RIV



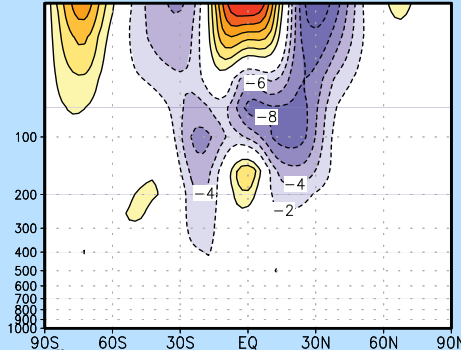
c) nRIV



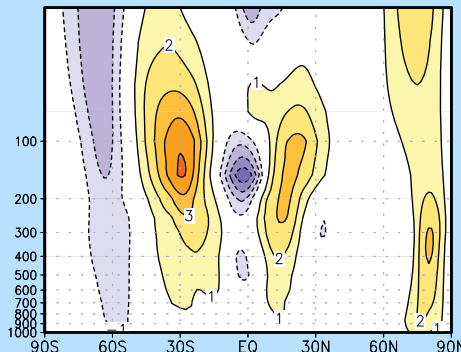
d) RIV - R2



e) nRIV - R2



f) RIV - nRIV



Zonal mean zonal wind

nRIV & RIV : satisfactorily reproduced.

separation between the polar night
and subtropical jet streams

nRIV run

underestimated

over the tropics and mid-latitude region

overestimated

over the tropics and polar region

RIV run

increased

over the mid-latitude region

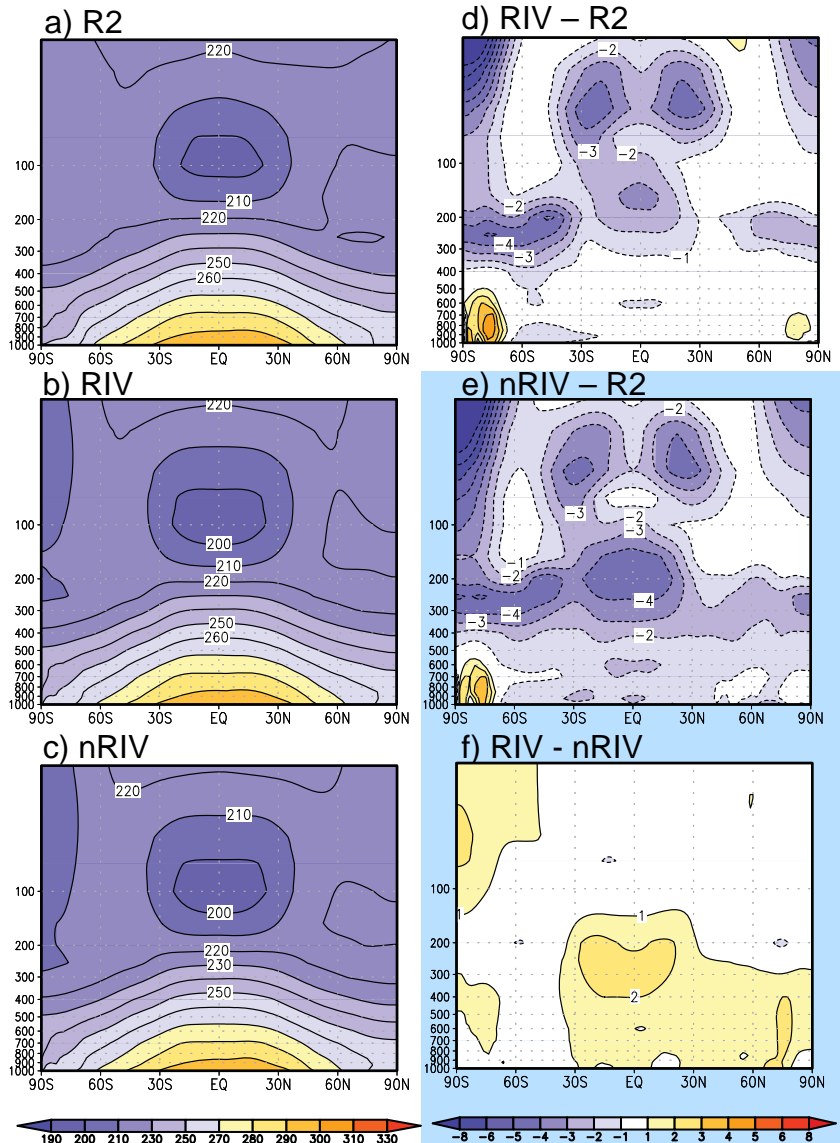
decreased

over the tropics

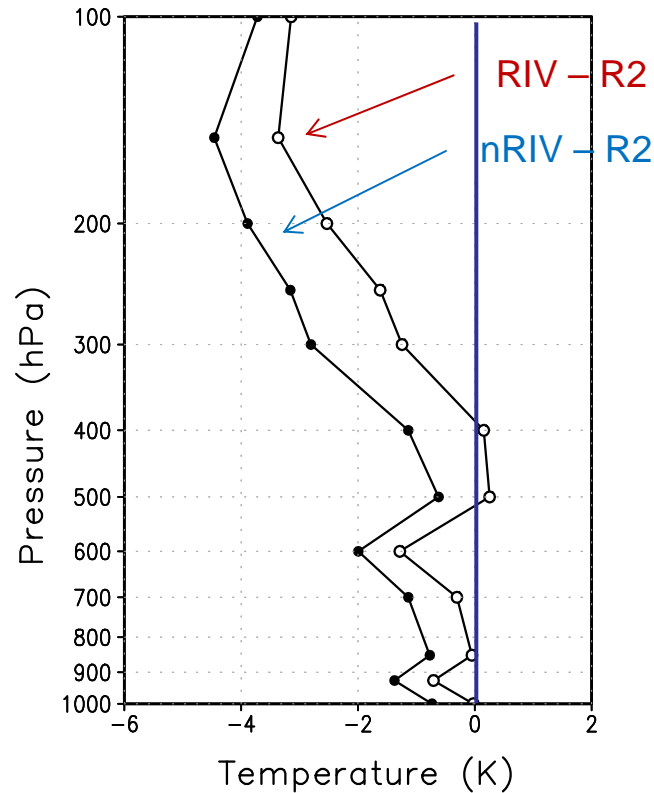
Effects of river routing processes



Zonal mean temperature



Vertical temperature profile



nRIV:

cold biases in most of the layer

RIV:

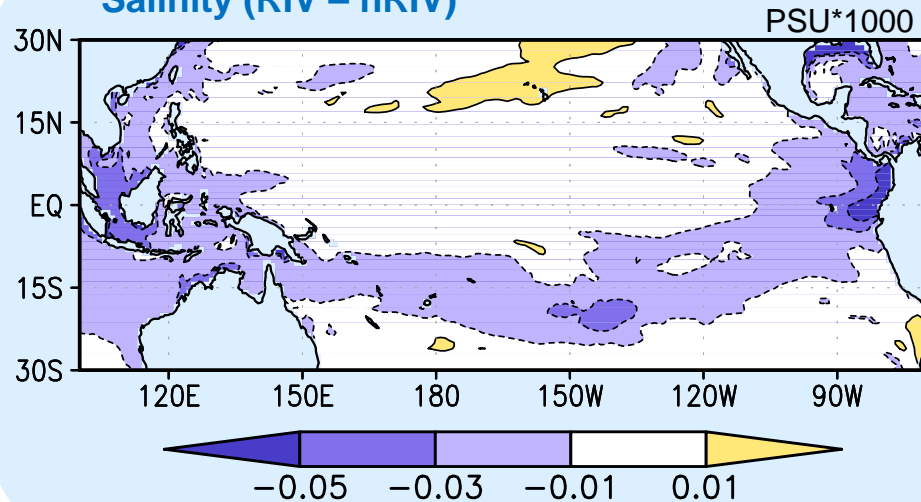
warmer than the nRIV experiment



Tropical Pacific Climate



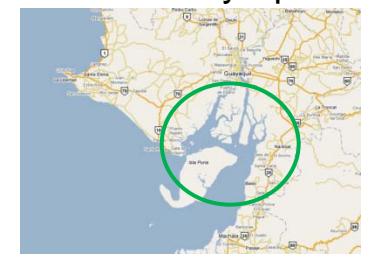
Salinity (RIV – nRIV)



Indonesia

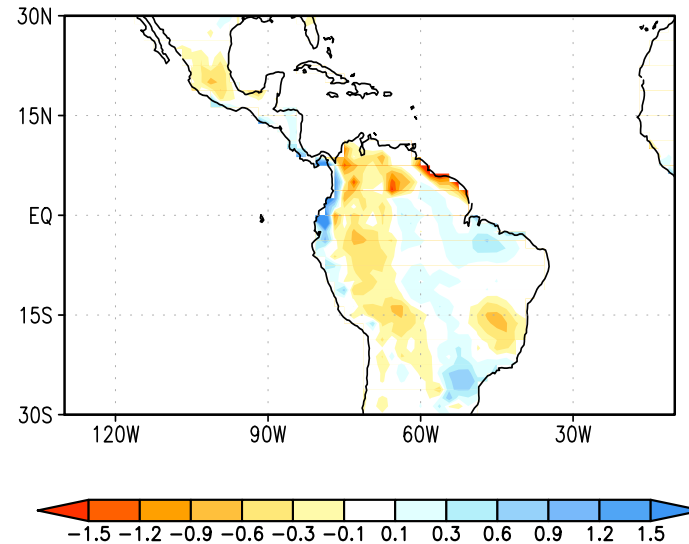
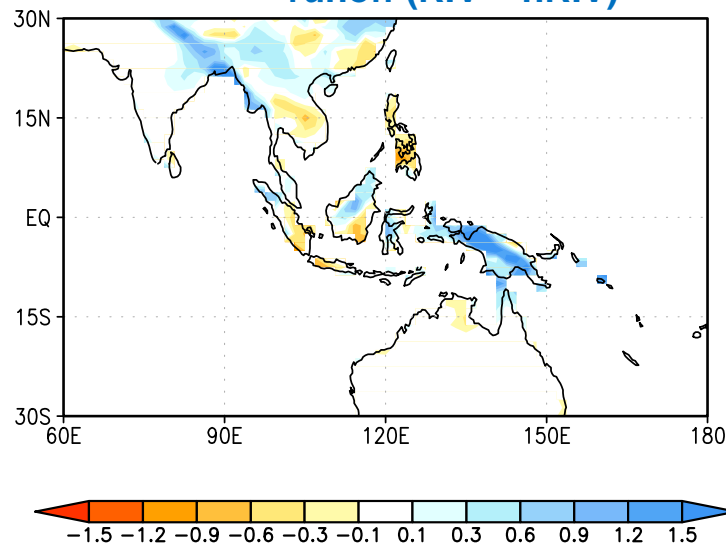


Gulf of Guayaquil

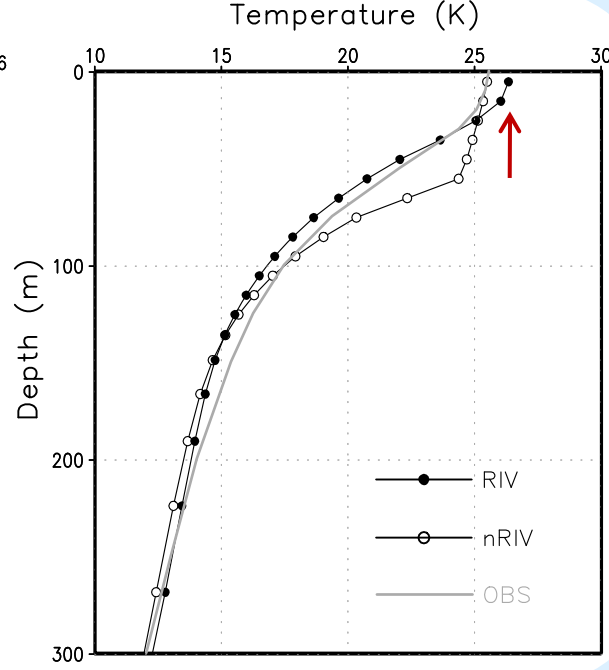
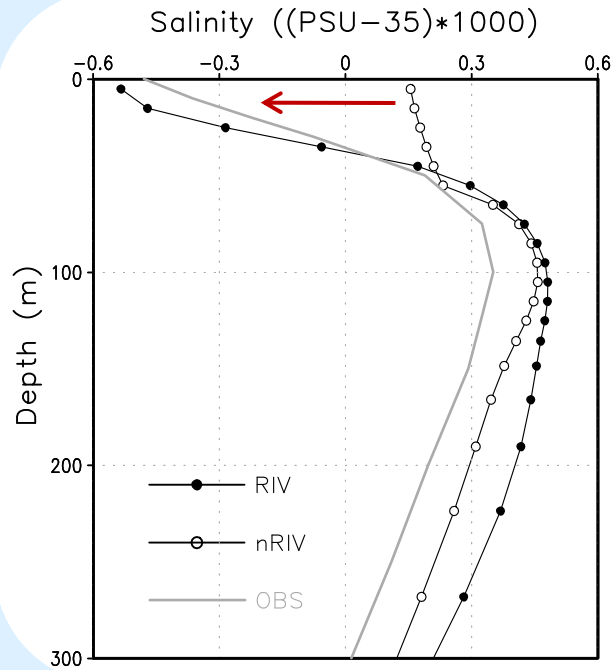


It is noticeable that **salinity** from the RIV run is **reduced** over the coastal region!!
→ It is due to the fact that **runoff is increased** (induced-precipitation, river discharge)

Surface + Subsurface
runoff (RIV – nRIV)



Tropical Pacific Climate



averaged for
eastern Pacific region
15° S ~ 15° N;
70° W ~ 110° W

—●— RIV
—○— nRIV
— OBS
(Levitus, 1998)

Surface salinity from the RIV run : **reduced!!**

→ due to **freshwater forcing**

(induced-precipitation, river discharge)

→ It is **closer to Levitus observed data**, although the results in deep ocean is not good.

Surface temperature from the RIV run: **increased!!**

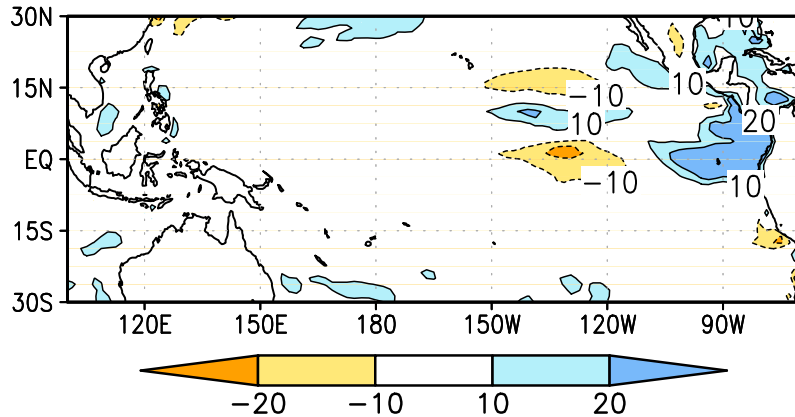
→ due to **shallower ocean mixed layer depth**; → leads to a **weakening of trade winds**

→ leads to a **reduction of upwelling**

→ produces an **increase in SST**

Tropical Pacific Climate

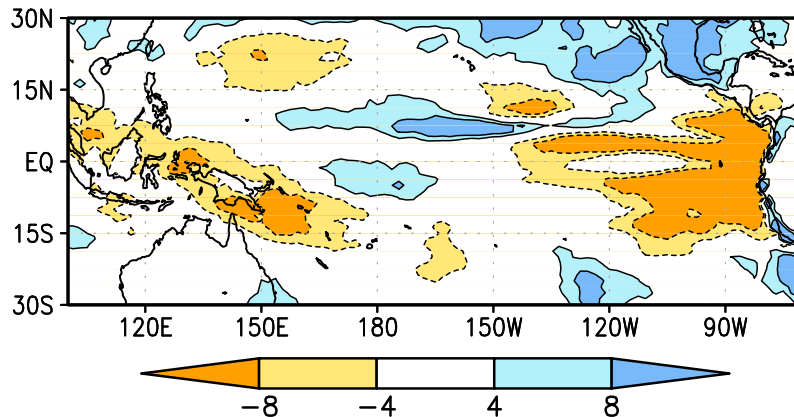
Latent heat flux (RIV – nRIV)



LH

increased over the eastern Pacific
due to warm SST and convection

Surface downward Shortwave flux (RIV – nRIV)



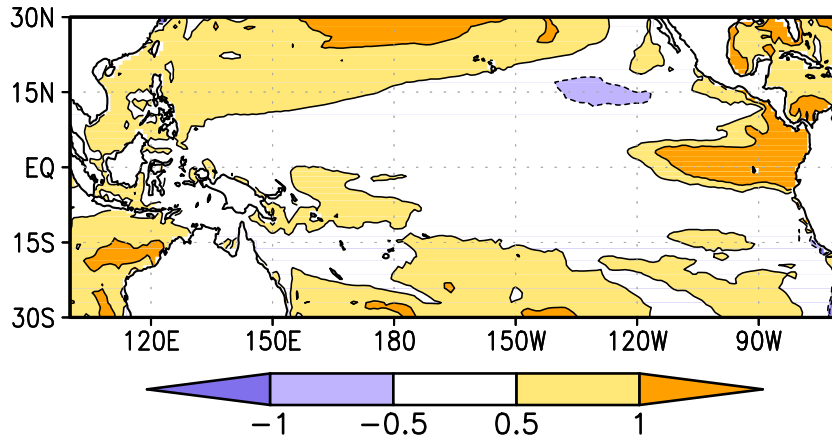
Surface downward SW flux

increased
over the mid-latitude region
decreased
over the tropics

→ It is closely affected by the reflection and absorption of solar radiation associated with the increase of cloud amount.

Tropical Pacific Climate

SST (RIV – nRIV)



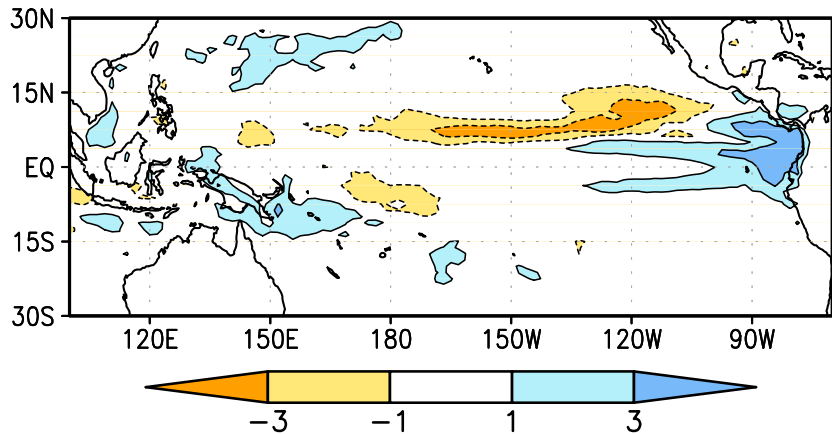
SST

increased over the eastern Pacific
due to reduced salinity

Precipitation

decreased
over the north of the equator in the central
Pacific,
increased
over the eastern equatorial Pacific

Precipitation (RIV – nRIV)



In other words,

: reduction of salinity → surface warming over the
eastern Pacific → increase in clouds by convection →
enhanced LH, reduced SW
→ increase in precipitation

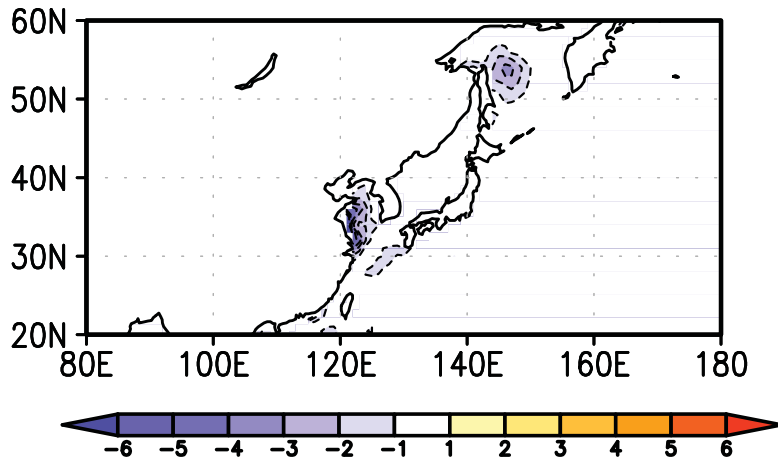
→ An overall increase of precipitation activity on a
globe reduces the biases in the large-scale features by
warming and moistening the troposphere.



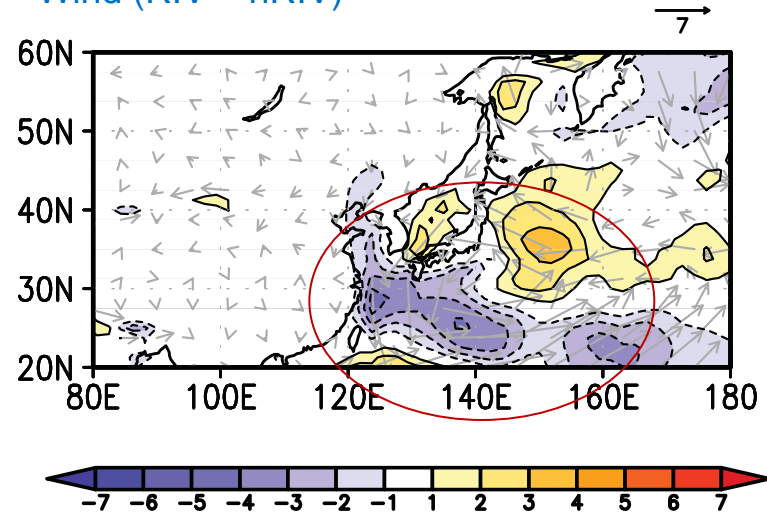
East Asian Monsoon (Summer:JJA)

Freshwater forcing also affects the EASM

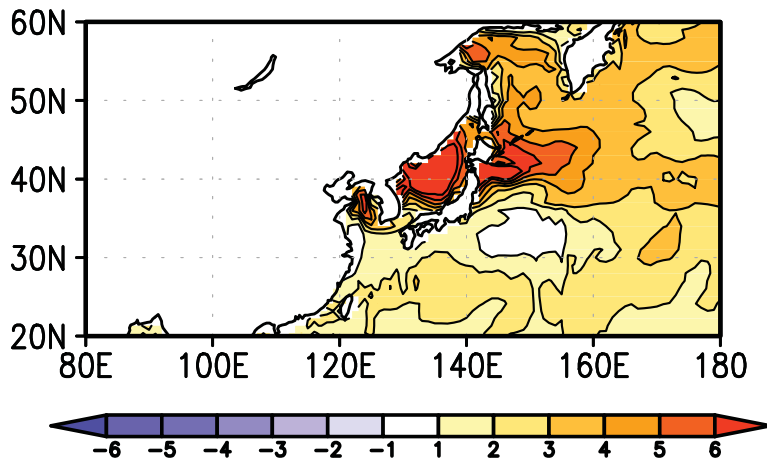
Salinity (RIV – nRIV)



Wind (RIV – nRIV)



Sea Surface potential Temperature (RIV – nRIV)



Reduced salinity

- increase in SST due to reduction of vertical mixing
- Alters the largescale circulation
- leads to a decrease in North Pacific high

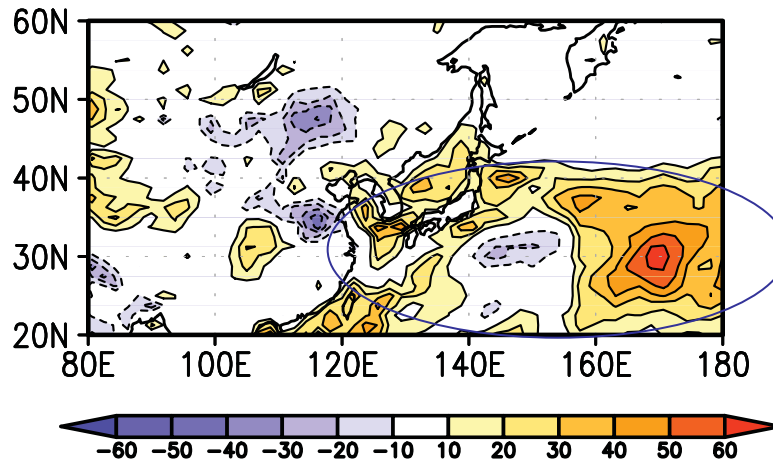
Cyclonic circulation

→ North Pacific High decrease



East Asian Monsoon (Summer:JJA)

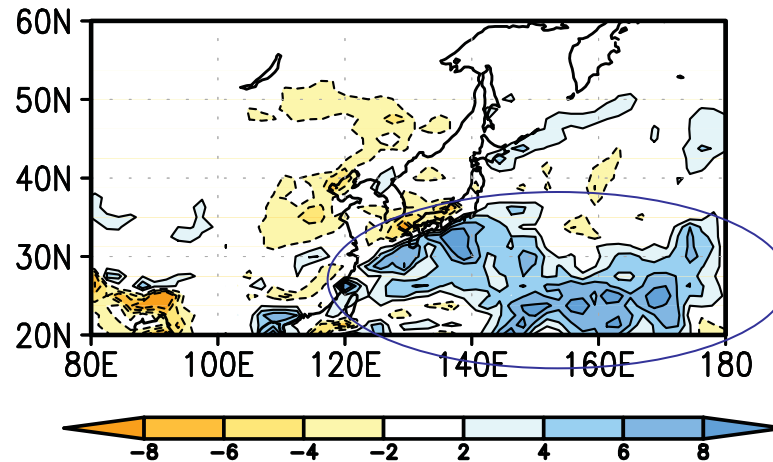
Latent heatflux (RIV – nRIV)



Cyclonic circulation
→ North Pacific High decrease



Precipitation (RIV – nRIV)



As a results

Latent heat increase
Precipitation increase

→ the **River Routing processes**
can affect the
East-Asian Summer Monsoon system



Summary



- ❖ The simulated **fresh water flux** into the oceans alters their salinity and may **affect the thermohaline circulation**.
- ❖ Recently, climate modeling community has focused on the **need for** routing models to track the **flow of water from continents to oceans at global scale**.
- ❖ The effect of river routing processes in a HadGEM2-AO on the simulated climate is significant. It is evident that **biases** in temperature and precipitation **are reduced**.
- ❖ In tropical Pacific, Sea Surface Temperature is increased by **reduction of trade winds** and **shallower mixed layer depth**, in turn, **precipitation is increase** due to convection.
- ❖ In East Asia, freshwater can affect **large-scale circulation**, as well as precipitation.



Conclusion



- ❖ River routing processes affect both ocean and large-scale circulation. Thus, it is **not negligible in climate simulation** since it alters the SST, which is the external boundary condition for the atmospheric model.
- ❖ There are a few issues to be cleared.
 - First, the structure of thermohaline needs to be verified over an observation.
 - Further, the modulation of surface evaporation due to the runoff over land needs to be considered.
- ❖ Despite these uncertainties, **the freshwater runoff** and its interaction with atmosphere **should be considered** in climate simulation



**Thank you
for your attention ~ !!!**